

# **Investigation for ground water Potentials Ataplot in Ifiteawka, Awkasouth local government area, Anambra state, Nigeria.**

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## **ABSTRACT;**

*The geophysical investigation report is the result of field electromagnetics survey using ADMT300s for the sinking of a successful borehole at a plot in Ifite Awka, Awka South Local Government Area, Anambra State. Pre-drilling geophysical surveys happen to be a scientific method of determining the existence of groundwater. This therefore, necessitated the commissioning of consulting hydro geologists to carry out this survey. The field electromagnetics survey using ADMT300s was therefore carried out at the proposed site to determine its water bearing formations. This report involves studies of the physiography, local geology and hydrogeology as well as the conduct of field electromagnetics survey using ADMT300s for the subsurface groundwater investigation. Based on the geological and geophysical studies, a borehole depth of not less than 823 feet (270 meters) is recommended to be drilled at the project site Ifite Awka, Awka South Local Government Area, Anambra.*

**Key words; Groundwater, Hydro geologist, Geophysical survey, Subsurface, River basin**

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## **I. INTRODUCTION**

In the face of increasing scarcity of water resources, there is a need for a holistic study of groundwater resources and potential of any proposed borehole drilling. The study area falls within Anambra River Basin. The most important aquiferous formation in the basin is the Ajali Formation. This formation is underlain by about 585m of the Mamu formation and Nkporo shale. A geophysical survey for groundwater is a cost-effective and non-intrusive method used to detect and assess potential sources of groundwater. It involves techniques such as electrical resistivity imaging and seismic methods to analyze subsurface geology and identify water-bearing formations. These surveys help in pinpointing the location of underground layers that contain water, which is critical for determining the best drilling locations for boreholes.

## **PROJECT LOCATION**

Ifite Awka is located at Awka South Government Area, Anambra state. Globally, it is at Latitude  $06^{\circ}14'22.8''$  North of the Equator and Longitude  $006^{\circ}05'42.3''$  East of Greenwich Meridian at an elevation of 91 meters.

## **GEOLOGY OF THE AREA**

Ifite Awka and environs including the project location lies at the crest of structure—the Awka–Orluridge, which acts as a groundwater divide. This structure is underlain by the Imo Clay/Shale geological unit. It is of Paleocene in age. The Imo Clay/Shale is a narrow stripe geological unit running roughly north–south with three sandy members, namely—the Ugwuoba Sandstone, The Imo Clay/Shale developed as thick bluish to greyish clays with a maximum thickness of about 1000m. This formation is essentially a aquiclude, except for the small lenticular sands, which constitute good aquifers in some localities. Entire formation was deposited in a marine environment during the Paleocene of the Cretaceous era.

## **HYDROGEOLOGY**

Hydro

geologically, Anambra falls within the Mamu River Basin. The River Mamu is a very important tributary of the Anambra River. It rises principally from the Eziana/Akokwah highland to the South, and is joined by numerous tributaries rising from the Enugu Ukwu/Agulu/Ekwulobia ridge to the east. The river flows roughly south–north, making a north west bend around Ebenebeto to join the Anambra river.

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The most important aquifers in Mamu River Basin are the Ajali Formation, the Uguwoaba and Amenyi Sandstone of the Imo Clay/Shale and the Nanka Sands. These aquiferous units however did not extend to the project location.

## II. METHODOLOGY

### GEOPHYSICAL INVESTIGATION

Groundwater behaves like a weak electrolyte. During the process of infiltration of runoff from rainwater into the subsurface where it is stored as groundwater, it passes through different rock columns, leaching some of its mineral content as it percolates downwards. This dissolved mineral content in groundwater makes it like a weak electrolyte. As a result, electrical geophysical approach is very suitable for groundwater surveys. Data from instrument capable of detecting electrical responses in the subsurface where utilised. The instrument and its measuring technique are here discussed.

### ADMT300S

This ADMT series instrument uses electromagnetic field of the earth as a working field source to examine the electrical structure of the subsurface. The electromagnetic response of the earth can reveal the electrical variations of the geological bodies at different depths in the subsurface via different frequency sampling. The depth of the electrical variations caused by various geologic conditions like electrolytic groundwater can be determined by sampling from higher frequency values to the lower frequency. Frequency is inversely proportional to wavelength which relates directly to depth. The electromagnetic wave propagation theory explained by Helmholtz equations shows that the propagation of electromagnetic wave passing through a filter (earth) obeys the Maxwell equation. Assuming an non-magnetic and uniformly conductive subterranean geological medium where there is no charge accumulation, then the Maxwell equation can be simplified to:

$$\left. \begin{aligned} \nabla^2 H + k^2 H &= 0 \\ \nabla^2 E + k^2 E &= 0 \end{aligned} \right\} \quad (1)$$

$H$  = Magnetic component  
 $E$  = Electrical component

Where  $k$  is called the wavenumber (or propagation coefficient) and it is expressed as:

$$k = [\omega^2 \mu \epsilon - i \omega \sigma \mu]^{1/2} \quad (2)$$

Considering that propagation coefficient,  $K$ , as a complex number, which implies that  $K$  has both the real and the imaginary part,  $K = b + ia$ ; where  $a$  is called the phase coefficient whereas  $b$  is the absorption coefficient. In the electromagnetic frequency range measured by this ADMT series range from 0.1 Hz to 5 kHz which are very low and as a result the displacement current caused. This further simplifies the  $K$  as:

(3)

A magnetic field with a change in Helmholtz equation induces a change in the electric field and as a result; we can have the magnetic to electric relationship:

$$\frac{E}{H} = \frac{-i \omega \rho}{k} \quad (4)$$

The surface impedance  $Z$  is defined as the ratio of the surface electric field and the horizontal component of the magnetic field. In the case of a uniform earth, this impedance is independent of the polarization of the incident field, and it related to the earth's resistivity and the:

$$Z = \frac{E}{H} = \sqrt{\omega \mu \rho e^{i\pi/4}} \quad (5)$$

The formula in equation 5 can be used to determine the resistivity of the earth layer:

$$\rho = \frac{1}{5f} \left| \frac{E}{H} \right|^2 \quad (6)$$

For the depth to the earth layer with the resistivity described in equation 6, the Skin depth approximation can be utilised on the basis that the media is non-magnetic:

$$\delta \approx 503 \sqrt{\frac{\rho}{f}} \quad (7)$$

Given the mathematical relationship in equation 7, the depth of penetration of electromagnetic wave is dependent on frequency and resistivity of the earth layer. The higher the resistivity, the greater the depth of penetration. The lower the frequency, also, the greater the penetration depth.

In ADMT 300s, a maximum depth of 300m is sampled and the lower resistivity interval can be suggestive of an aquifer.

**Table 1. PROBE LOCATIONS**

S/N	PROBE ID	LATITUDE (°) (Degree decimals)	LONGITUDE (°) (Degree decimals)	ELEVATION (m)	Date
1	FID60	N06°14'22.8 <sup>II</sup>	E007°05'42.3 <sup>II</sup>	91	January, 17 <sup>th</sup> 2025.
2	FID61	N06°14'22.0 <sup>II</sup>	E007°05'40.3 <sup>II</sup>	97	January, 17 <sup>th</sup> 2025.

### III. RESULTS AND DISCUSSION

2N0 probe points were acquired (Table 1) at the project site and are denoted as **FID60** and **FID61**. The tomographic section is shown in Figure 1 & 2. A depth of 300 meters was achieved at each probe point.

#### **FID60**

This was carried out within the project site about 10 meters from the entrance gate. 2N0 ground water horizons were detected. The first ground water horizon occurred at a depth range of 160 meters. The second ground water horizon occurred at a depth range of 200-

250 meters. From the 2D contour map section (Figure 1), the subsurface is characterized by high resistivity (impermeable) rock types, colour coded red with intercalations of wet and saturated zones colour coded light green and blue respectively. Underlying rock units are dominantly **Shale/Clay** with interbeds of **Mudstone/Siltstone/Sandstone**. Saturation connectivity largely do exist, hence may enhance conductivity.

#### **FID61**

This was carried out still within the project site at about 28 meters from the entrance gate. Ground water horizons were detected at a depth of 165 meters.

However, deeper abstraction may be possible at the project site at a depth range of 250 meters to 300 meters. From the 2D contour map section, the subsurface is characterized by high resistivity (impermeable) rock types, colour coded red with intercalations of wet zones colour coded light green. Underlying rock units are dominantly **Shale/Clay** with interbeds of **Mudstone/Siltstone/Sandstone** hence is dicey to recommend this point.

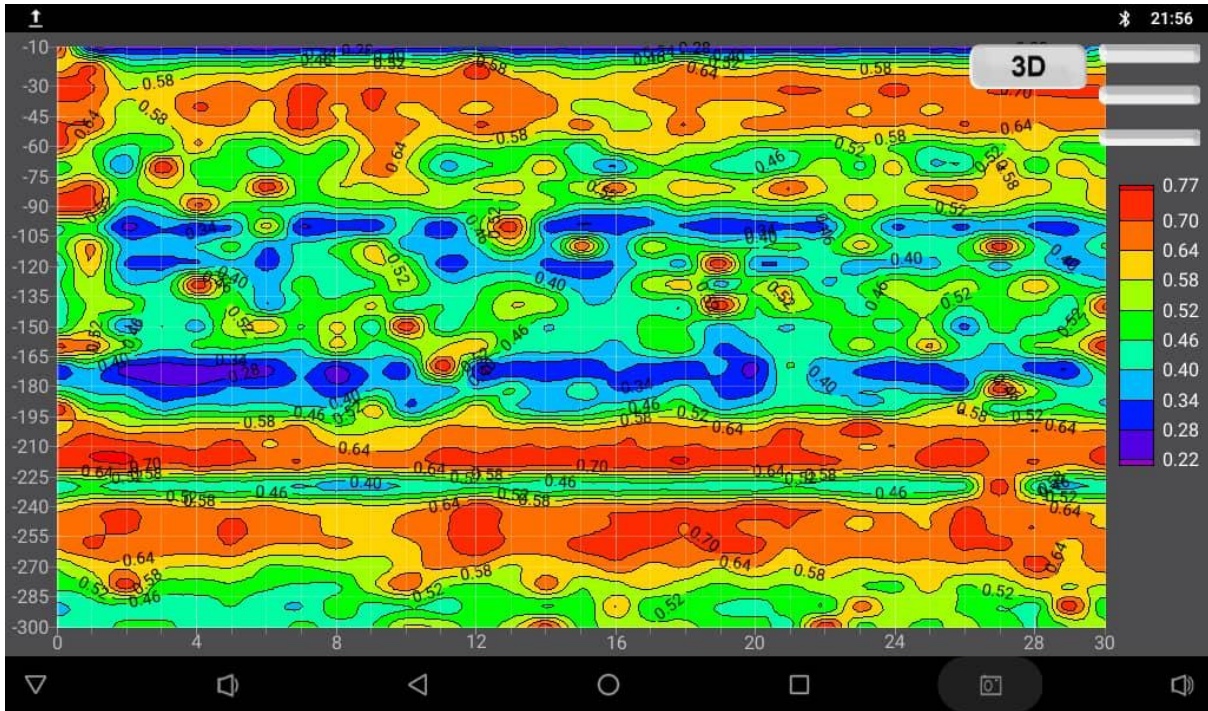


Figure 1:FID60

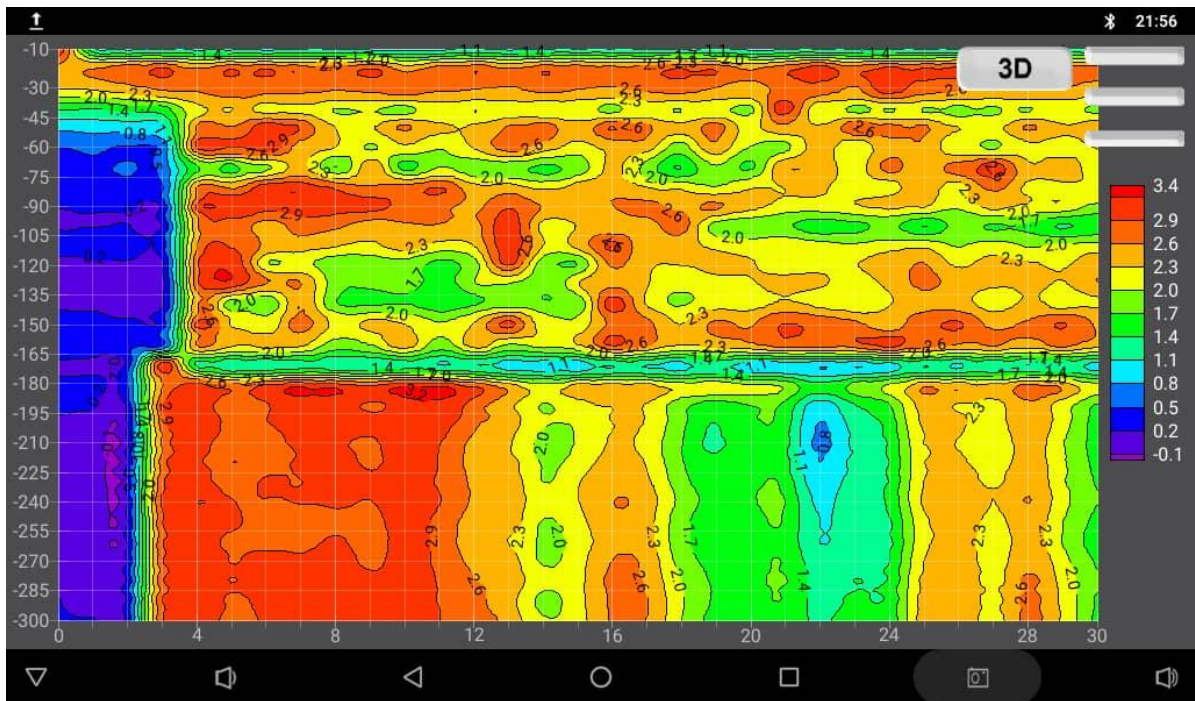


Figure 2:FID61

#### IV. RECOMMENDATIONS

Based on the geological and geophysical studies, a borehole depth of not less than **823 feet (270 meters)** is recommended to be drilled at the project site Ifite Awka, Awka South Local Government Area, Anambra State.

Geophysical logging of the boreholes should be undertaken before insertion of screens and casings.

The entire drilling programme should be supervised by a competent and experienced Geologist/Water Engineer.

The drilling method recommended is the direct rotary method, using a **Competent Rotary Rig**.

Adequate preparations should be made by provision of adequate quantity of chemical to forestall the incidence of collapse, since the formation may be mostly sandy.

#### REFERENCE

- [1]. Barlow, P. M., &Leake, S. A. (2012). **Groundwater resources and their management in the United States**. Environmental Earth Sciences, 66(3), 847-857. <https://doi.org/10.1007/s12665-011-1526-1>
- [2]. Bear, J. (2018). **Hydraulics of groundwater**. Dover Publications.
- [3]. Carrillo-Rivera, J. J., & Peralta, M. G. (2010). **Hydrogeological investigations in semi-arid regions**. Hydrogeology Journal, 18(5), 1181-1194. <https://doi.org/10.1007/s10040-010-0616-z>
- [4]. Domenico, P. A., & Schwartz, F. W. (1998). **Physical and chemical hydrogeology** (2nd ed.). John Wiley & Sons.
- [5]. Fetter, C. W. (2017). **Applied hydrogeology** (5th ed.). Pearson.
- [6]. Freeze, R. A., & Cherry, J. A. (1979). **Groundwater**. Prentice-Hall.
- [7]. Healy, R. W., & Ronan, D. M. (2006). **Technical review of groundwater models used for hydrogeological investigations**. Journal of Hydrology, 324(1-4), 26-40. <https://doi.org/10.1016/j.jhydrol.2005.08.020>
- [8]. Harter, T., &Dahm, C. N. (2016). **Hydrogeological investigations in agricultural regions**. Agricultural Water Management, 176, 12-27. <https://doi.org/10.1016/j.agwat.2016.06.003>
- [9]. Lerner, D. N., & Harris, S. (2009). **Groundwater contamination and hydrogeological investigations**. Water Resources Research, 45(6), W06420. <https://doi.org/10.1029/2008WR007419>
- [10]. Mustafa, M., & Ali, R. (2021). **Integrated hydrogeological investigation of the aquifer system in the Khyber Pakhtunkhwa region of Pakistan**. Journal of Hydrology, 599, 126338. <https://doi.org/10.1016/j.jhydrol.2020.126338>
- [11]. Simmers, I., & Martens, J. (2001). **Hydrogeological investigations in arid zones: Conceptual approaches and techniques**. Environmental Geology, 40(5), 611-616. <https://doi.org/10.1007/s002540100282>
- [12]. Wood, W. W., & Sanford, W. E. (1995). **Groundwater investigation in arid regions: A practical guide**. Journal of Hydrology, 167(1-4), 121-135. [https://doi.org/10.1016/0022-1694\(94\)02699-8](https://doi.org/10.1016/0022-1694(94)02699-8)